



Norwegian  
Meteorological  
Institute

# Surface wave-induced drift

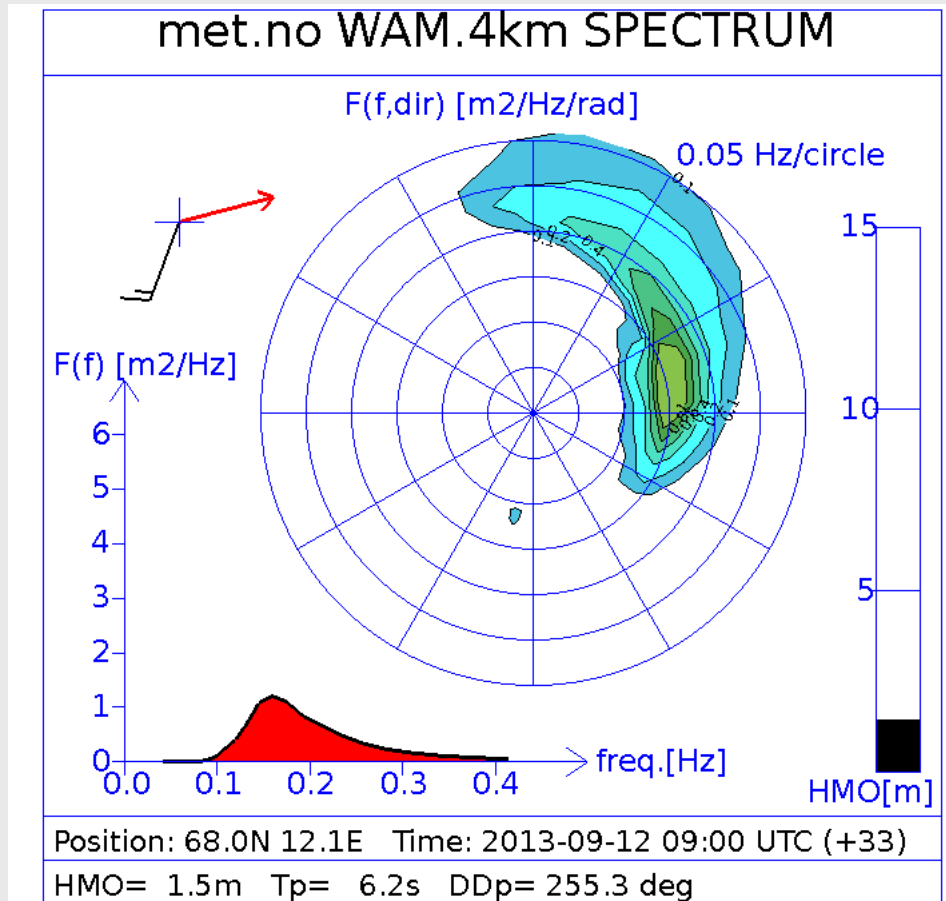
Kai H. Christensen

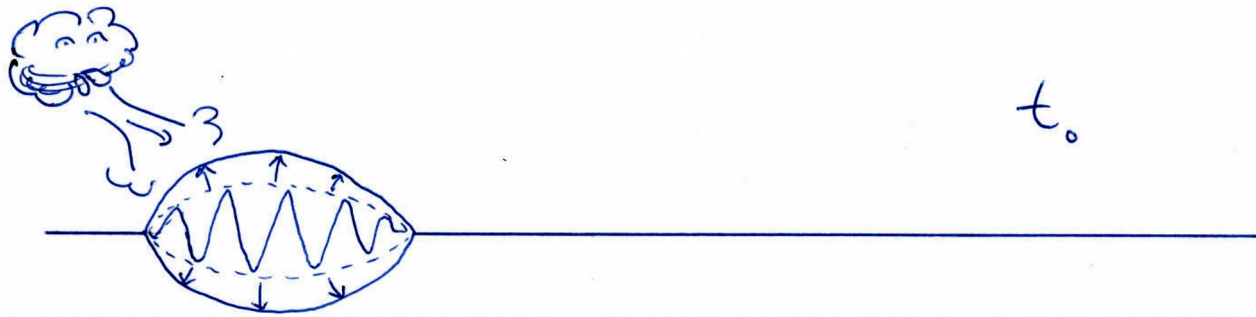
Research and Development Department

Division for Ocean and Ice

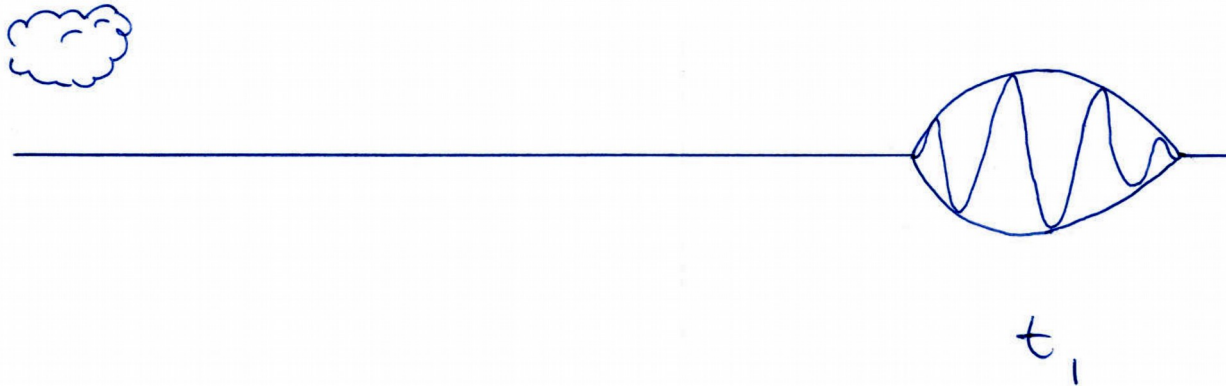
# The energy balance equation

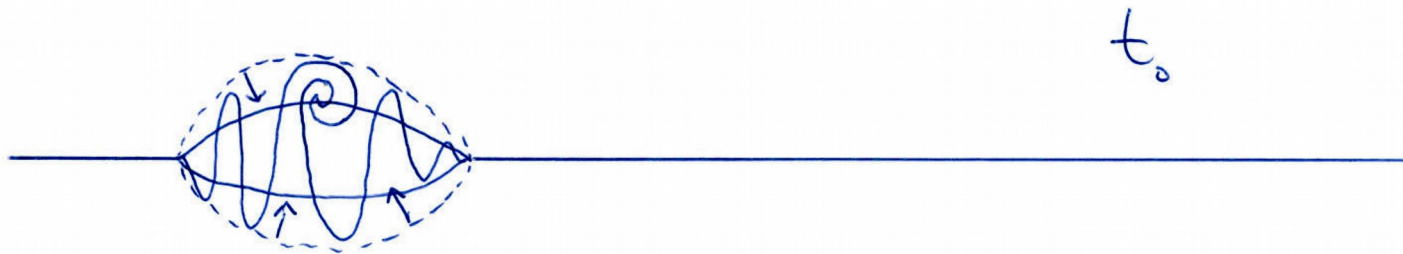
$$\frac{\partial F}{\partial t} + \nabla \cdot (\mathbf{c}_g F) = S_{input} + S_{dissip} + S_{nonlin}$$



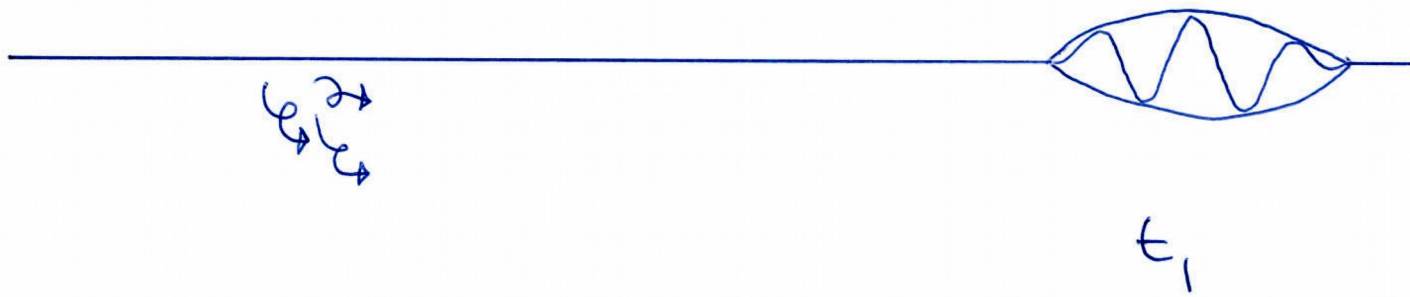


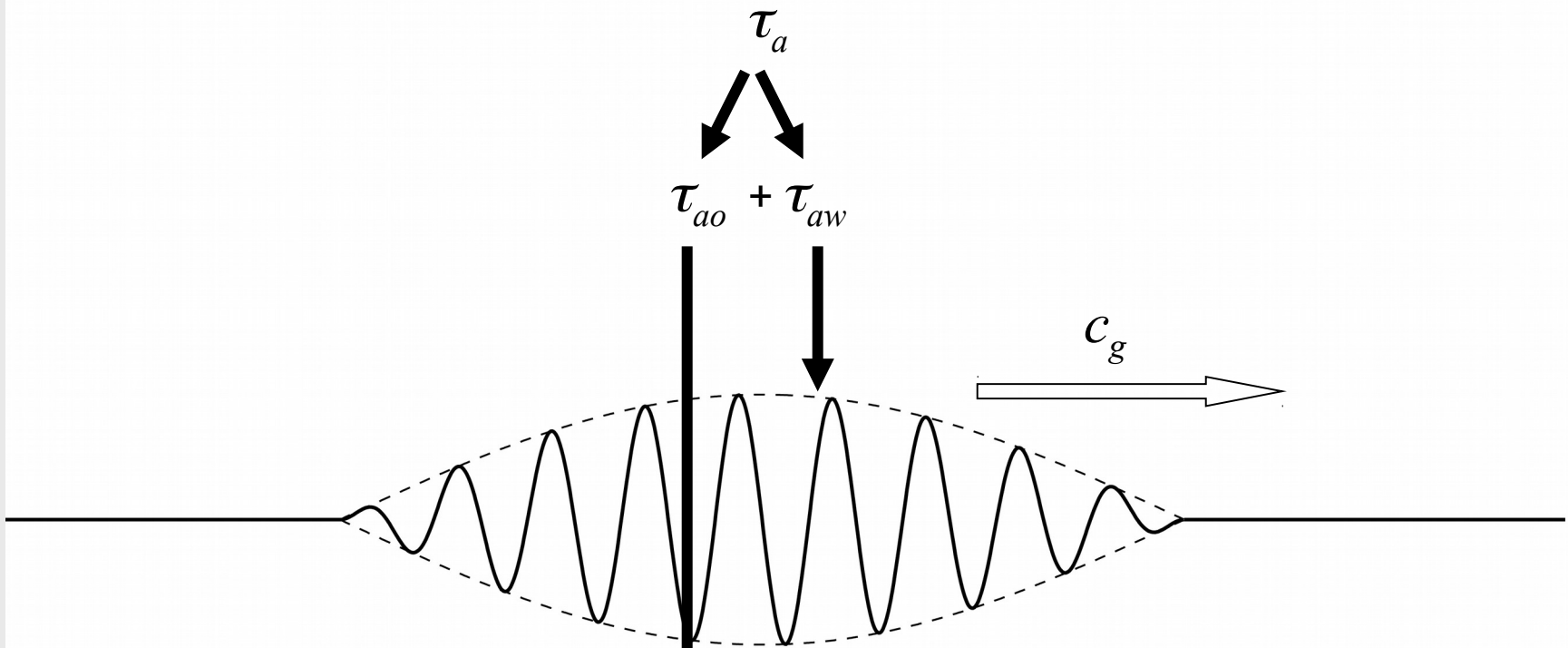
$C_g \Rightarrow$





$C_g$  





Weber et al (2006):

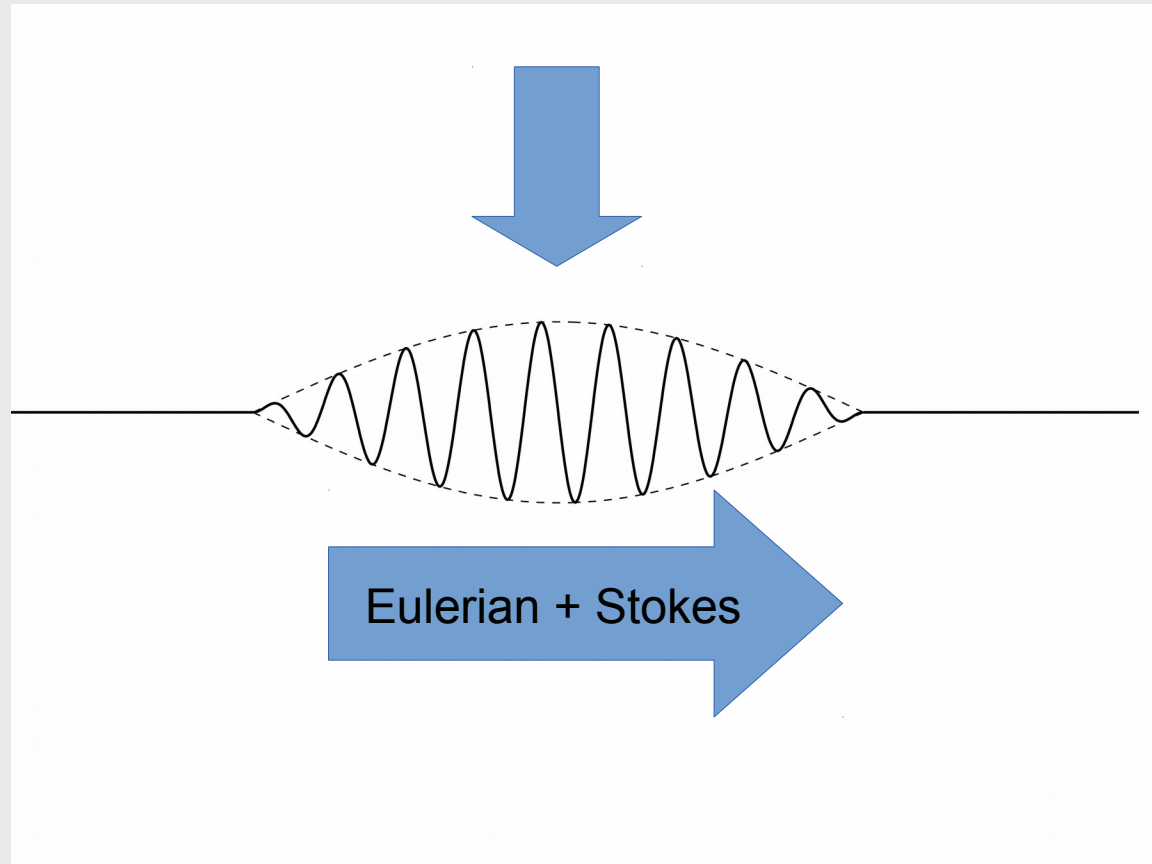
$$U_s = \int_{-D}^{\xi} u_s dz$$

$$\frac{\partial U_s}{\partial t} + \nabla \cdot (c_g U_s) = \frac{\tau_{aw}}{\rho} - \frac{\tau_{wo}}{\rho}$$

LOCAL FORCING  
VS  
ADVECTION BY GROUP

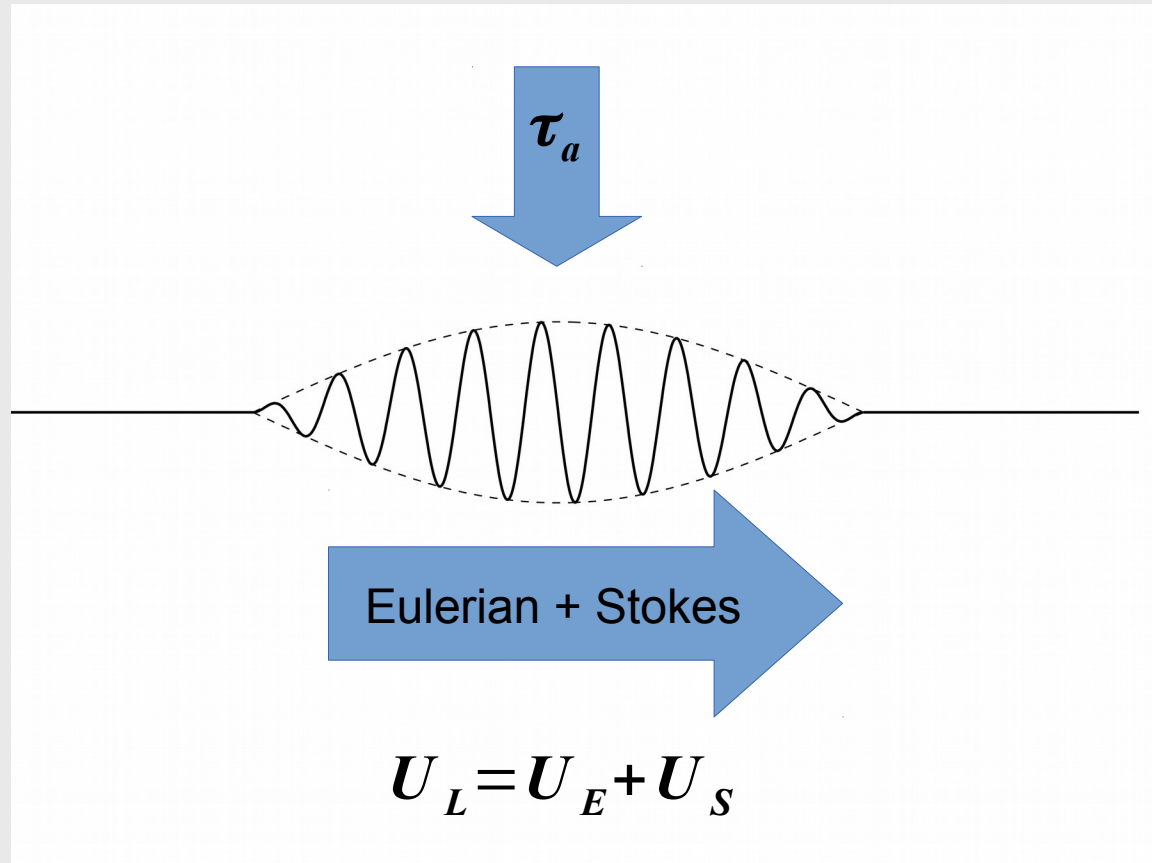
# Ekman dynamics with waves

Momentum flux from atmosphere



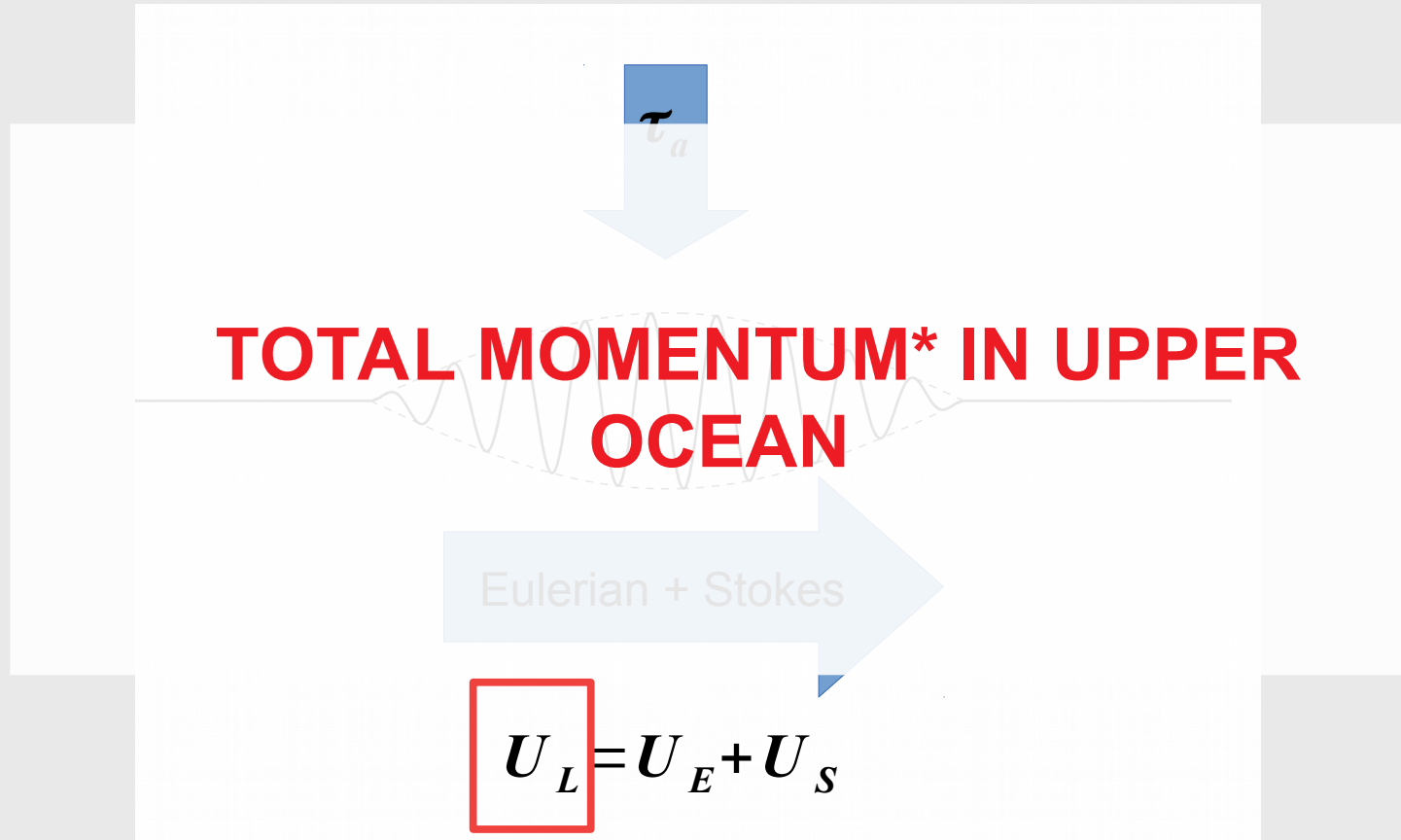
# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$



# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

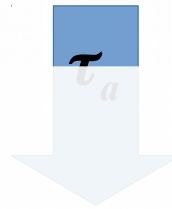


\* Per unit density

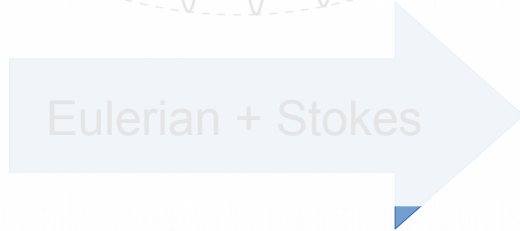


# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$



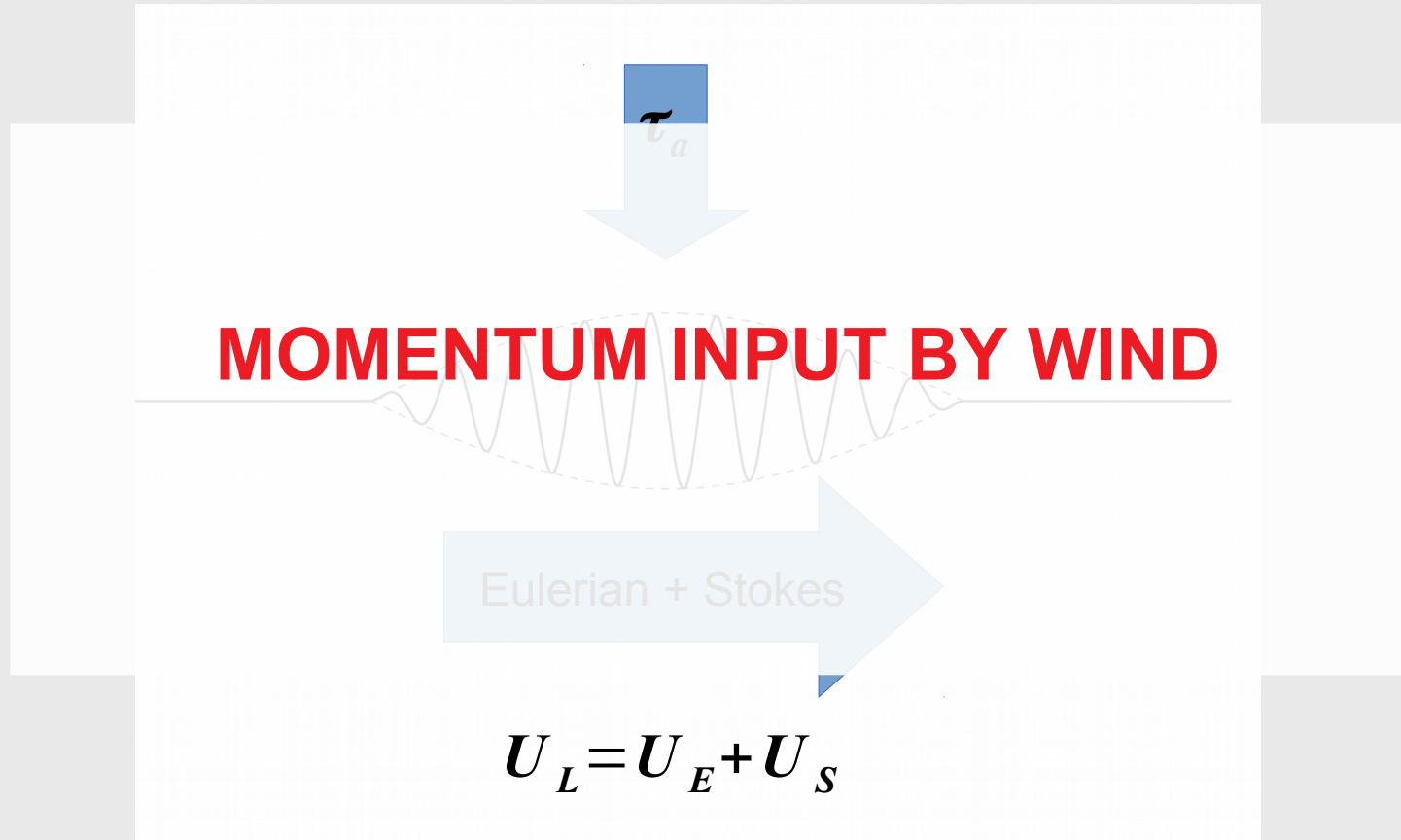
**ROTATING REFERENCE FRAME**



$$\mathbf{U}_L = \mathbf{U}_E + \mathbf{U}_S$$

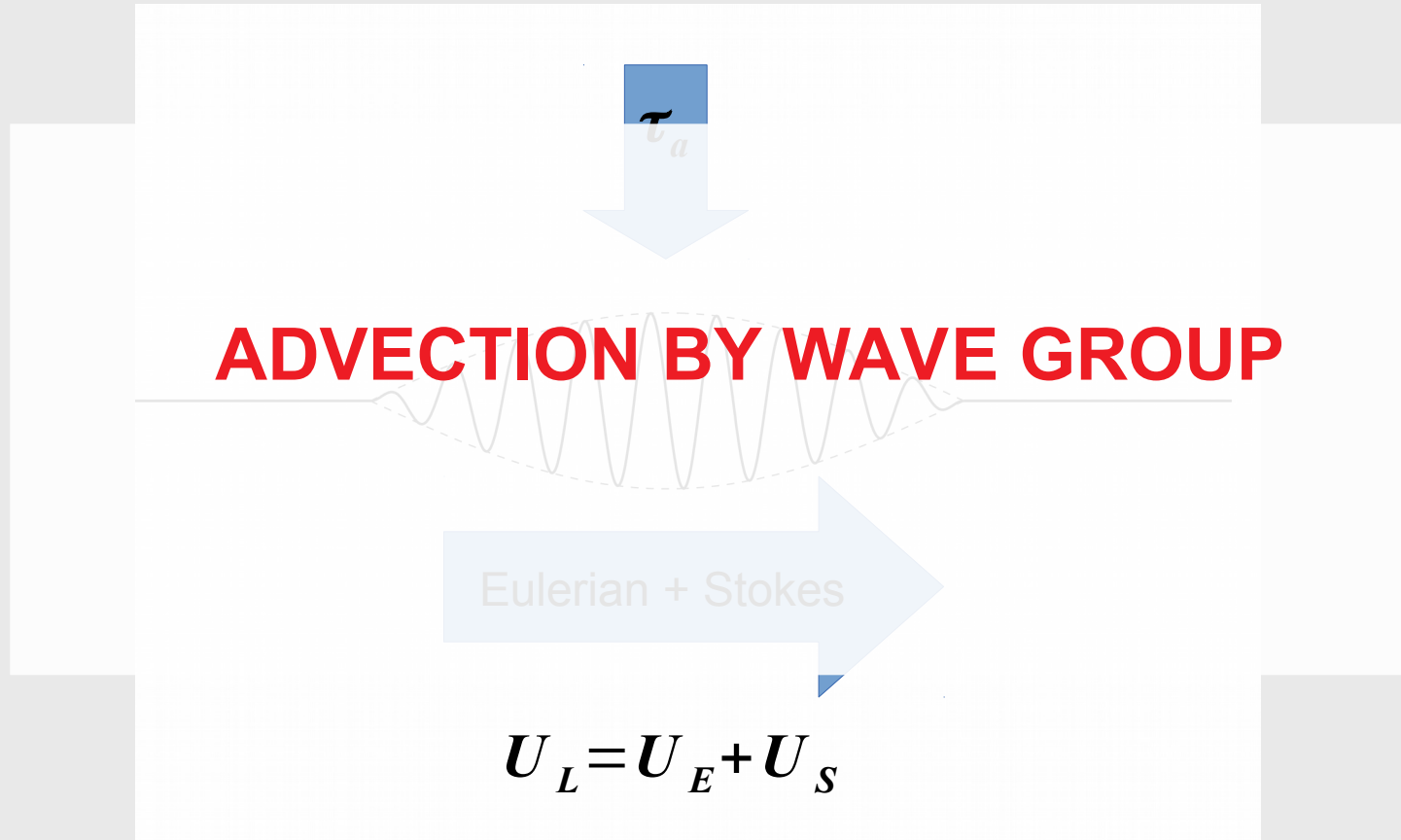
# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$



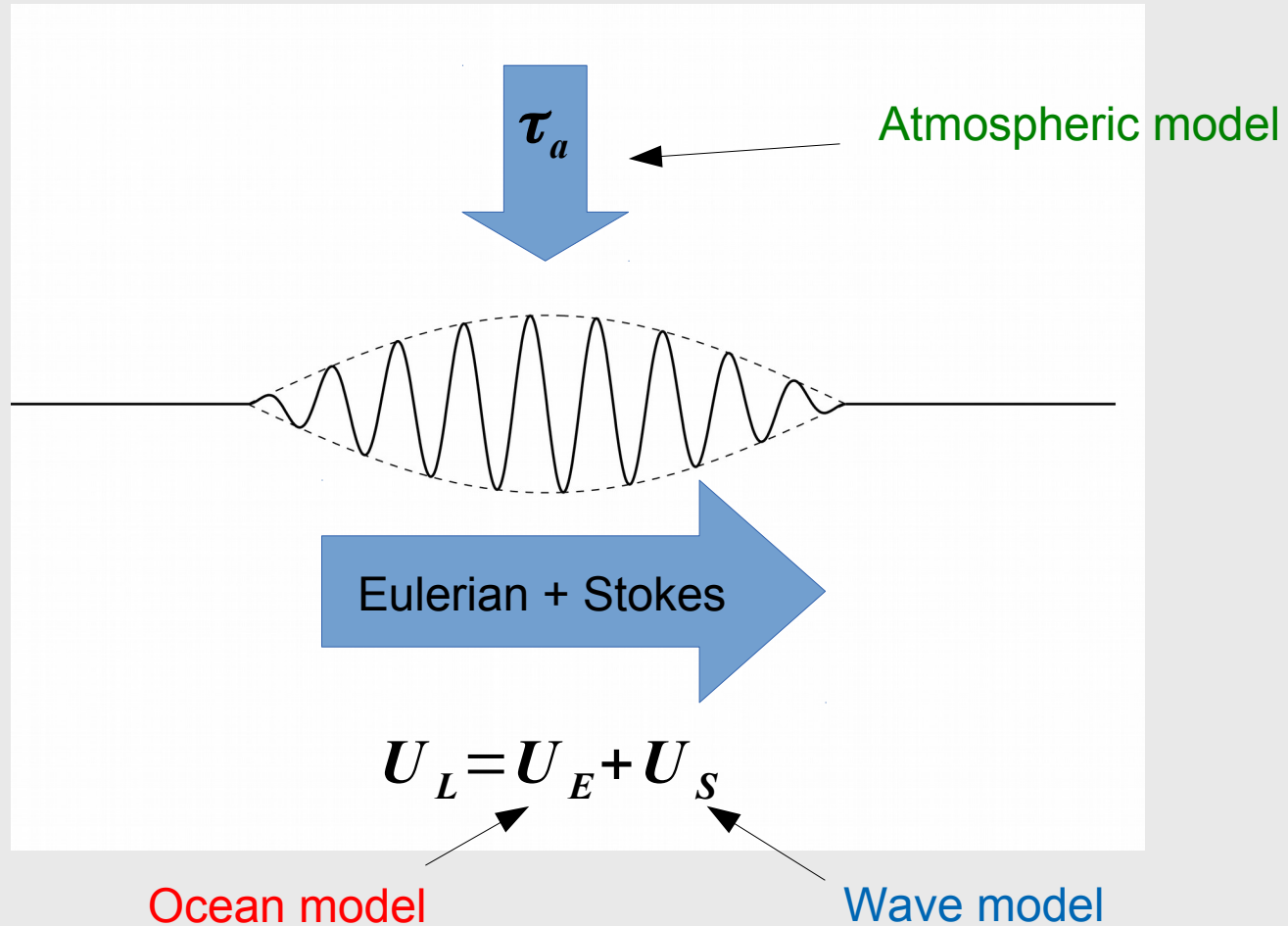
# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$



# Ekman dynamics with waves

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$



$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(expanding)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(expanding)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(rewriting the «wave model»)

$$\frac{\partial \mathbf{U}_S}{\partial t} + \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{aw}}{\rho} - \frac{\boldsymbol{\tau}_{wo}}{\rho} \quad \longrightarrow \quad \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{wo}}{\rho} + \frac{\partial \mathbf{U}_S}{\partial t}$$

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(expanding)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(rewriting the «wave model»)

$$\frac{\partial \mathbf{U}_S}{\partial t} + \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{aw}}{\rho} - \frac{\boldsymbol{\tau}_{wo}}{\rho} \quad \longrightarrow \quad \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{wo}}{\rho} + \frac{\partial \mathbf{U}_S}{\partial t}$$

(substitute)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{wo}}{\rho} + \frac{\partial \mathbf{U}_S}{\partial t}$$

$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(expanding)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S)$$

(rewriting the «wave model»)

$$\frac{\partial \mathbf{U}_S}{\partial t} + \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{aw}}{\rho} - \frac{\boldsymbol{\tau}_{wo}}{\rho} \quad \longrightarrow \quad \frac{\boldsymbol{\tau}_{aw}}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{wo}}{\rho} + \frac{\partial \mathbf{U}_S}{\partial t}$$

(substitute)

$$\frac{\partial \mathbf{U}_E}{\partial t} + \frac{\partial \mathbf{U}_S}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_{ao}}{\rho} + \frac{\boldsymbol{\tau}_{wo}}{\rho} + \frac{\partial \mathbf{U}_S}{\partial t}$$

(simplify and use definition for ocean stress)

$$\frac{\partial \mathbf{U}_E}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_S) = \frac{\boldsymbol{\tau}_o}{\rho}$$



$$\frac{\partial \mathbf{U}_L}{\partial t} + f \mathbf{k} \times \mathbf{U}_L = \frac{\boldsymbol{\tau}_a}{\rho} - \nabla \cdot (\mathbf{c}_g \mathbf{U}_s)$$

Summary: We take the momentum budget for the **mean flow and the waves combined**, and then remove the parts specific to the «wave model». We are then left with the **«ocean model»**.

Coriolis-Stokes force

$$\frac{\partial \mathbf{U}_E}{\partial t} + f \mathbf{k} \times (\mathbf{U}_E + \mathbf{U}_s) = \frac{\boldsymbol{\tau}_o}{\rho}$$

Sea-state dependent momentum flux (surface stress)

# Tying it all together

$$\frac{\boldsymbol{\tau}_o}{\rho} = \frac{\boldsymbol{\tau}_a}{\rho} - \left[ \frac{\partial \mathbf{U}_s}{\partial t} + \nabla \cdot (\mathbf{c}_g \mathbf{U}_s) \right]$$

$$\frac{\partial F}{\partial t} + \nabla \cdot (\mathbf{c}_g F) = S_{input} + S_{dissip} + S_{nonlin}$$

$$\mathbf{U}_s = g \frac{F}{c}$$

$$\boldsymbol{\tau}_o = \boldsymbol{\tau}_a - \rho g \int_0^\infty \int_0^{2\pi} \frac{\mathbf{k}}{\omega} (S_{input} + S_{dissip}) d\theta df$$

# Notes

- Conceptual model!
  - Horizontal gradients are ignored
  - Finite depth effects ignored
  - Variable density is ignored
  - Vertical variation not described
  - Coupling ocean -> wave ignored
  - ...
- Spectral resolution of wave model can be an issue
- Similar coupling between the waves and the atmosphere, important process for weather forecasting